Temporal and spatial variation in active layer depth in the McMurdo Sound Region, Antarctica

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Abstract: A soil climate monitoring network, consisting of seven automated weather stations, was established between 1999 and 2003, ranging from Minna Bluff to Granite Harbour and from near sea level to about 1700 m on the edge of the polar plateau. Active layer depth was calculated for each site for eight successive summers from 1999/2000 to 2006/2007. The active layer depth varied from year to year and was deepest in the warm summer of 2001–02 at all recording sites. No trends of overall increase or decrease in active layer depth were evident across the up-to-eight years of data investigated. Average active layer depth decreased with increasing latitude from Granite Harbour (77°S, active layer depth of 90 cm) to Minna Bluff (78.5°S, active layer depth of 22 ± 0.4 cm), and decreased with increasing altitude from Marble Point (50 m altitude, active layer depth of 49 ± 9 cm) through to Mount Fleming (1700 m altitude, active layer depth of 6 ± 2 cm). When all data from the sites were grouped together and used to predict active layer depth the mean summer air temperature, mean winter air temperature, total summer solar radiation and mean summer wind speed explained 73% of the variation (R2 = 0.73).

Received 9 December 2008, accepted 8 June 2009

Key words: active layer prediction, altitudinal gradient, Cryosol, Gelisol, latitudinal gradient, permafrost, soil temperature

Introduction

Global temperatures are predicted by the International Panel for Climate Change (IPCC 2007, table SPM.1) to increase by 1.8–4.0°C by 2099. The polar regions are expected to have a magnified response to the projected increase in global temperatures compared to more temperate regions (Kane et al. 1991). Increased air temperature will have an impact on the active layer and permafrost. The active layer is defined as the layer of ground subject to annual freezing and thawing cycles in areas underlain by permafrost (Linell & Tedrow 1981, Guglielmin 2006), while permafrost is defined as the ground that remains continuously at or below 0°C for more than two years (Muller 1947, Linell & Tedrow 1981).

The active layer has been predicted to be significantly impacted by the projected increase in air temperature over the next several decades (Kane et al. 1991) particularly in Alaska where the mean annual air temperature is about 0°C. Kane et al. (1991) predicted that the active layer depth in Alaska will increase by 22 cm if mean annual air temperature warms by 2°C over the next 50 years, and an increase of 43 cm depth was predicted if air temperature was to warm by 4°C over a 50 year period.

The active layer depth in Antarctic soils provides a potential indicator of climate change (Guglielmin 2004) due to its sensitivity to changes in air temperature (Kane et al. 1991, Conovitz et al. 2006). Vegetation, which influences the relationship between active layer depth and surface climate parameters, is absent in the Ross Sea region of Antarctica (Guglielmin 2006).

The active layer depth in Antarctica has been shown to be spatially variable (Campbell et al. 1998, Bockheim & Hall 2002, Bockheim et al. 2007, Ikard et al. 2009) and is also dependent on the climatic zone in which the site is located (Weyant 1966, Campbell & Claridge 1987). For example, Coastal Zone sites may have active layer depths of about 50–60 cm (Campbell et al. 1998, Bockheim et al. 2007) whereas sites in the Inland Mountain subzone may have active layer depths on the order of about 5–10 cm (McKay et al. 1998, Bockheim et al. 2007).

With projected increases in global air temperature potentially impacting the active layer and permafrost, it is important to understand how the active layer varies interannually, what drives the interannual variation and the trends in active layer depth across latitude and altitude, before attempting to predict the impacts of climate change on active layer depth. Measurements of the active layer from a single season in the McMurdo Sound region of Antarctica have been conducted previously (e.g. Guglielmin et al. 2003, Conovitz et al. 2006) but there is a paucity of work concerning the interannual variation in active layer depth.

Seven automated weather stations have been established in the McMurdo Sound Region since 1999 and are located at Wright Valley at the base of Bull Pass (established
Marble Point (1999), Scott Base (1999), Victoria Valley (1999), Mount Fleming (2002), Granite Harbour (2003), and Minna Bluff (2003) (Fig. 1). Each station records data every hour, including air temperature, solar radiation, and soil temperature, at a range of depths.

The objectives of this paper are: to determine the active layer depth, in each summer, at each of the seven sites; to examine how the active layer depth varies between summers, with latitude, and with increasing altitude; and to determine the influence of climatic factors including air temperatures and solar radiation on active layer depth.

### Site descriptions, instrumentation, and methods

#### Site descriptions

The automated weather stations (Fig. 1) provide a latitudinal transect, near sea level, from Minna Bluff through to Granite Harbour, including the Scott Base and Marble Point stations. Complementary to the latitudinal component, the stations established at Marble Point through Wright Valley, Victoria Valley and Mount Fleming provide an altitudinal transect from 50 to 1700 m above sea level (Table I). Detailed soil descriptions, and analytical data are available at [http://soils.usda.gov/survey/scan/antarctica/index.html](http://soils.usda.gov/survey/scan/antarctica/index.html). All sites have no vegetation and are occasionally covered by a thin layer of snow, and are subject to a range of forcing factors such as winter storms, wind, and temperature. The sites are positioned in the McMurdo Sound region, from 70 to 1700 m above sea level, with a latitudinal range from 77°31'32.4''S to 25°10.6''S, and an altitudinal range from 50 to 1700 m above sea level.

#### Table I. Site location and descriptions of automated weather stations established since 1999 in the McMurdo Sound region, Antarctica.

<table>
<thead>
<tr>
<th>Site</th>
<th>Location</th>
<th>Elevation</th>
<th>aspect, Mean annual air temp.</th>
<th>Soil parent material</th>
<th>General comments</th>
</tr>
</thead>
<tbody>
<tr>
<td>Minna Bluff</td>
<td>78°30'41.6''S N side of Minna Bluff Peninsula, 75 km W of Scott Base buildings</td>
<td>163 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>163 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>163 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>163 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
</tr>
<tr>
<td>Scott Base</td>
<td>77°30'53.6''S Scott Base buildings</td>
<td>100 m, N 38 m, SSE, slope of 6</td>
<td>100 m, N 38 m, SSE, slope of 6</td>
<td>100 m, N 38 m, SSE, slope of 6</td>
<td>100 m, N 38 m, SSE, slope of 6</td>
</tr>
<tr>
<td>Marble Point</td>
<td>77°24'2.7''S 400 m E of the Wilson–Patterson Glacier</td>
<td>50 m, flat, weakly patterned ground</td>
<td>50 m, flat, weakly patterned ground</td>
<td>50 m, flat, weakly patterned ground</td>
<td>50 m, flat, weakly patterned ground</td>
</tr>
<tr>
<td>Granite Harbour</td>
<td>77°00'23.7''S 10 m wide bouldery cliff and sea ice edge immediately to the south</td>
<td>4 m, N, gently sloping</td>
<td>4 m, N, gently sloping</td>
<td>4 m, N, gently sloping</td>
<td>4 m, N, gently sloping</td>
</tr>
<tr>
<td>Wright Valley</td>
<td>76°51'57''E Wright Valley near the base of Bull Pass</td>
<td>150 m, S, gently sloping</td>
<td>150 m, S, gently sloping</td>
<td>150 m, S, gently sloping</td>
<td>150 m, S, gently sloping</td>
</tr>
<tr>
<td>Victoria Valley</td>
<td>77°31'32.4''S Valley floor</td>
<td>30 m 410 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>30 m 410 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>30 m 410 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>30 m 410 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
</tr>
<tr>
<td>Mount Fleming</td>
<td>77°32'42.7''S Mount Fleming, near edge of polar plateau</td>
<td>1700 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>1700 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>1700 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
<td>1700 m, flat, strongly developed gravelly sand over gravelly clay loam</td>
</tr>
</tbody>
</table>

*Mean annual air temperature could only be calculated for Mount Fleming in 2007, as this was the only year in which a full dataset was available.

### The objectives of this paper

To determine the influence of climate factors including air temperatures and soil radiation on active layer depth, to determine how the active layer depth varies between summers, with latitude, and with increasing altitude, and to determine the influence of climatic factors including air temperatures and solar radiation on active layer depth at each of the seven sites. The objectives of this paper are to determine the active layer depth, in each summer, at each of the seven sites; to examine how the active layer depth varies between summers, with latitude, and with increasing altitude; and to determine the influence of climatic factors including air temperatures and solar radiation on active layer depth.
average of 36 days per summer at Scott Base, 13 days per summer at Marble Point and one hour per summer at Wright Valley (Wall et al. 2004). Liquid moisture rarely occurs at Mount Fleming due to cold temperatures, Minna Bluff would probably have a similar moisture regime to Scott Base, and Victoria Valley is expected to be similar to Wright Valley. Granite Harbour has meltwater from an adjacent hillside flowing through the subsurface soil for much of the summer period.

Surface albedos are 10% at Scott Base, 25% at Marble Point, and 21% at Wright Valley (Balks et al. 2002). Given the similarity of geological materials, the albedo at Minna Bluff and Mount Fleming are expected to be similar to the albedo at Scott Base. Victoria Valley and Granite Harbour would be expected to have albedos similar to Wright Valley.

The apparent thermal diffusivities of soils at Scott base and in the McMurdo Dry Valleys were reported to be relatively low (1.2 x 10^{-7}–1.4 x 10^{-7} m^2 s^{-1}) and thermal conductivities were similar for all soils (about 0.2 Wm^{-1} K^{-1}) and relatively constant through the range of soil moisture content found in the field (Campbell et al. 1997). Ikard et al. (2009) reported similar mean bulk apparent thermal diffusivities (2.9 x 10^{-7} –116 x 10^{-7} m^2 s^{-1}) in the Taylor Valley with the higher values in moist soils near lake margins.

**Instrumentation**

Soil pits were dug into the ice cement, where present, to depths of 120 cm, or as deep as possible, to install the temperature probes. Care was taken to ensure the soil was returned as closely as possible to the original site condition.

Soil temperatures used to calculate active layer depth in this study were measured using two kinds of sensors. MRC (Measurement Research Corporation, Gig Harbor, WA) temperature probes were installed to 120 cm depth at Marble Point, Wright Valley and Victoria Valley, and to 115.6 cm depth at Scott Base. Each MRC probe measures temperatures at 11 depths over the 120 cm length of the probe. Temperature sensors (model 107; Campbell Scientific, Logan, UT) were used at Mount Fleming (to 75 cm depth) and Minna Bluff (to 84 cm depth up until the 2006 summer). A combination of MRC and 107 temperature probe measurements were used at Granite Harbour (to 90 cm depth) and at Minna Bluff (to 112 cm depth) after 2006. The location and depths of all temperature probes at each site are available at http://soils.usda.gov/survey/scan/antarctica/index.html.

Incoming solar radiation was measured 3 m above ground level using a pyranometer (LiCor model LI200X; Campbell Scientific, Logan, UT) at Wright Valley, Scott Base, Marble Point, Minna Bluff, Granite Harbour and Mount Fleming, and 2 m above ground level at Victoria Valley.

Air temperature was measured using an RM Young RTD temperature probe (model 43347; Campbell Scientific, Logan, UT) at Wright Valley (1.6 m above ground level).

![Fig. 2. Maximum summer active layer depths at seven sites in the McMurdo Sound region, from December 1999 to January 2007.](image1)

A Vaisala temperature and relative humidity probe was used at Scott Base (model HMP35C until January 2005, when it was replaced by HMP45C; Campbell Scientific, Logan, UT) (1.6 m), Victoria Valley (HMP45C) (2 m), and at Marble Point (HMP45C) (1.6 m). Air temperature at Minna Bluff was measured using a Vaisala temperature and relative humidity probe (HMP35C) until January 2006, after which a Campbell Scientific 107 probe in a solar radiation shield was installed 2 m above ground level. A Campbell Scientific 107 temperature probe in a solar radiation shield was used at Mount Fleming (1.6 m) and at Granite Harbour (2 m).

![Fig. 3. Mean summer (December–January) air temperature (°C) at all sites in McMurdo Sound region of Antarctica.](image2)
Table II. Date (d/m/y) when the maximum active layer depth occurred in each summer at seven sites in the McMurdo Sound region. NE denotes that the site was not established in this summer. The date of maximum depth was determined when the soil temperature reached its maximum at the two depths either side of the 0°C isotherm.

<table>
<thead>
<tr>
<th>Summer</th>
<th>Minna Bluff</th>
<th>Scott Base</th>
<th>Marble Point</th>
<th>Granite Harbour*</th>
<th>Wright Valley</th>
<th>Victoria Valley</th>
<th>Mount Fleming</th>
</tr>
</thead>
<tbody>
<tr>
<td>1999–00</td>
<td>NE</td>
<td>13/01/00</td>
<td>17/01/00</td>
<td>NE</td>
<td>18/01/00</td>
<td>17/01/00</td>
<td>NE</td>
</tr>
<tr>
<td>2000–01</td>
<td>NE</td>
<td>19/01/01</td>
<td>30/01/01</td>
<td>NE</td>
<td>26/01/01</td>
<td>24/01/01</td>
<td>NE</td>
</tr>
<tr>
<td>2001–02</td>
<td>NE</td>
<td>16/01/02</td>
<td>14/01/02</td>
<td>NE</td>
<td>18/01/02</td>
<td>08/01/02</td>
<td>NE</td>
</tr>
<tr>
<td>2002–03</td>
<td>NE</td>
<td>13/01/03</td>
<td>29/01/03</td>
<td>NE</td>
<td>15/01/03</td>
<td>05/01/03</td>
<td>24/01/03</td>
</tr>
<tr>
<td>2003–04</td>
<td>17/01/04</td>
<td>12/01/04</td>
<td>13/01/04</td>
<td></td>
<td>10/01/04</td>
<td>10/01/04</td>
<td>24/01/04</td>
</tr>
<tr>
<td>2004–05</td>
<td>23/01/05</td>
<td>14/01/05</td>
<td>22/01/05</td>
<td>12/01/05</td>
<td>20/01/05</td>
<td>18/01/05</td>
<td>19/12/04</td>
</tr>
<tr>
<td>2005–06</td>
<td>09/01/06</td>
<td>04/01/06</td>
<td>18/01/06</td>
<td>18/01/06</td>
<td>18/01/06</td>
<td>18/01/06</td>
<td>31/12/05</td>
</tr>
<tr>
<td>2006–07</td>
<td>13/01/07</td>
<td>31/01/07</td>
<td>30/01/07</td>
<td>29/01/07</td>
<td>14/01/07</td>
<td>26/01/07</td>
<td>09/12/06</td>
</tr>
</tbody>
</table>

*Date of occurrence at Granite Harbour was taken at the date of the maximum soil temperature at the deepest sensor (90 cm).

All sensors were connected to dataloggers (CR10X-2M; Campbell Scientific, Logan, UT). Atmospheric measurements were taken once every 10 sec and soil measurements were taken once every 20 min. All measurements were averaged every hour and the hourly average recorded. Data were manually downloaded each January.

Active layer determination

Active layer depth was determined by the intercept of the annual maximum soil temperature profile with the 0°C isotherm (Burn 1998, Guglielmin et al. 2003). This method is reliable because it is replicable and not influenced by subjective observers (Burn 1998) and data at our sites are continuously recorded all year. In this study, the austral summer is defined as the time from 1 December–31 January. For calculation of active layer depth at least 7 and up to 13 temperature-depth measurements were used at each site.

Statistical analysis

Correlation and regression analyses were undertaken using Minitab (Minitab Inc 2007) and Microsoft Excel (v. 2007).

Results and discussion

Active layer depth

The maximum active layer depth at the seven sites (Fig. 2) varied spatially and temporally (between summers). The active layer depths at sites in the coastal climatic zone (Scott Base (mean 32 cm, SD 6), Marble Point (mean 49 cm, SD 9), and Granite Harbour (> 90 cm)) were generally deeper than those further from the open sea (Minna Bluff (mean 22 cm, SD 4), Wright Valley (mean 46 cm, SD 7), Victoria Valley (mean 21 cm, SD 4), and Mount Fleming (mean 6 cm, SD 2)).

The between-summer variability in active layer depths (Fig. 2) showed a similar pattern to mean summer air temperature (Fig. 3). While there were between-summer differences there was no trend of increasing or decreasing active layer depth over the period investigated. Due to the limited length of the available record and the interannual variability in active layer depth it will be necessary to continue the monitoring for a much longer period to determine any long-term trends in active layer depth.

While Mount Fleming had the coldest summer temperatures and the shallowest active layer, the pattern was not consistent. Granite Harbour, with the deepest active layer, had mean summer air temperatures that were not significantly different from those at the other low altitude sites (Adlam 2009). While Wright Valley and Marble Point had similar active layer depths the mean summer air temperature was markedly warmer at Wright Valley.

Local micro-topography may influence active layer depth (Cannone et al. 2008). The greater active layer depth at Granite Harbour was probably the result of heat transfer from meltwater that was observed to percolate through the subsurface from the adjacent hillside. In the Taylor Valley deeper active layer depths have been shown to occur at moist sites, than in adjacent dry sites, as a result of increased thermal conductivity in moist soils (Ikard et al. 2009), and as a result of heat transfer via sub-surface water flow (Conovitz et al. 2006). The shallow active layer at Mount Fleming was within the range expected for sites in the Inland Mountain subzone (about 5–10 cm, McKay et al. 1998, Bockheim et al. 2007). The Mount Fleming site was near the edge of the polar plateau and subject to, near constant, strong winds derived from cold air drainage from the polar plateau. The Minna Bluff site, situated on the north side of the Minna Bluff Peninsula, was also a site subject to frequent severe winds. All other sites were on generally flat areas with minimal effects from meltwater runoff and had relatively “normal” Antarctic wind conditions.

The date of maximum soil temperatures used to calculate the active layer depth varied interannually (Table II). The time of maximum active layer depth ranged in occurrence from late December to late January.

Over half (57%) of the recorded dates of maximum thaw were between 10 and 20 January with a further quarter (27%) in late January (21–30). Early January (1–9) maximums were recorded twice at Victoria Valley and once each at Scott Base and Minna Bluff. At the relatively cold and windy...
Mount Fleming site, maximum temperatures were recorded in December in three summers with the remaining two summers recording maximum temperatures in late January. In the cold summer of 2000–01 three of the four recording sites had maximum temperatures in late January, probably due to the extended period of cool, cloudy, weather experienced across the region in the first half of January 2001. There were no obvious trends between date of maximum thaw and specific site factors, such as albedo or microtopography, given that all sites varied from summer to summer. The factor most likely to be affecting the timing of maximum thaw is the summer weather pattern. A two or three week pattern of warm sunny, or cool cloudy, weather in early January could have a strong influence.

Change in active layer depth with latitude and altitude

As would be widely expected the active layer depth showed clear latitudinal (Fig. 4) and altitudinal (Fig. 5) gradients. The latitudinal gradient in Fig. 4 may be somewhat exaggerated due to the local topographic effects at Granite Harbour, discussed previously.

Hatherton (1990) suggested that the maximum dry adiabatic lapse rate could be used to estimate temperature differences based on elevation differences in Antarctica, which in turn explains the decrease in active layer depth with increasing elevation (Campbell & Claridge 2006). Between the Wright Valley and Mount Fleming, the observed temperature increase had the same slope as that predicted by the dry adiabatic lapse rate (Fig. 6), with the decrease in air temperature conforming to that predicted due to the adiabatic lapse rate. However, the dry adiabatic lapse rate did not predict the temperature difference between Marble Point and the higher altitude, Wright Valley. The Marble Point site is within the Coastal Mountain climate zone of Campbell & Claridge (1987), while the other sites are within the Inland Mountain zone. It is clear that the differences in active layer depth between the climate zones, as illustrated by the difference between Marble Point and Wright Valley, are not attributable to altitude alone.

The 2001–02 warm summer

The active layer depths at all operational sites (Scott Base, Marble Point, Wright Valley and Victoria Valley) were deepest in the summer of 2001–02 which was one of the warmest summers on record in the McMurdo Dry Valleys (Foreman et al. 2004, Barrett et al. 2008) and also in the western Antarctic Peninsula and the Bellingshausen Sea (Massom et al. 2006). The warm temperatures experienced in the 2001–02 summer in the Antarctic Peninsula were
caused by the presence of a “blocking high” (Turner et al. 2002, Massom et al. 2006), that persisted from mid-September to February, and was linked, by the British Antarctic Survey, to the positive phase of the Southern Annular Mode (Massom et al. 2006). Turner et al. (2002) noted that there was an anomalously low mean sea level pressure situation over the same period (2001–02) in the Ross Ice Shelf region, which may have been a contributing factor to the warmer surface air temperatures and thus deeper active layer depths recorded.

**Relationship between active layer depth and climatic variables**

Mean summer air temperature and mean summer windspeed

Summer air temperature (December–January) has been documented to have a strong influence on the depth of the active layer (Zhang et al., 1997, Conovitz et al. 2006, Bockheim et al. 2007). Guglielmin (2004) noted the strong relationship between ground surface and air temperature \( R^2 > 0.9 \) in the Ross Sea Region. Conovitz et al. (2006) found that the active layer depth in the McMurdo Dry Valleys could respond to changes in air temperature within a time period as short as 24 hours.

Active layer depth correlated positively with mean summer air temperature (1 December–31 January) at Scott Base \( R = 0.87 \), Marble Point \( R = 0.91 \), Wright Valley \( R = 0.92 \) and Victoria Valley \( R = 0.94 \). However, Mount Fleming \( R = 0.24 \) and Minna Bluff \( R < 0.1 \) showed no relationship between mean summer air temperature and active layer depth. Both Minna Bluff and Mount Fleming are subject to strong winds (with annual average wind speed of 8.8 m s\(^{-1}\)) at Minna Bluff and 10.1 m s\(^{-1}\) at Mount Fleming) and had shorter temperature records available.

Mean summer wind speed gave a positive correlation with active layer depth between summers at Minna Bluff \( R = 0.93 \), Wright Valley \( R = 0.76 \), Victoria Valley \( R = 0.65 \) and Mount Fleming \( R = 0.56 \). However there was no relationship between mean summer wind speed and active layer depth at Marble Point \( R = -0.2 \) or Scott Base \( R = 0.2 \). Granite Harbour was not included as all values for active layer depth were greater than 90 cm and no correlation coefficient was able to be calculated.

Mean winter air temperature

Within each site over time, the correlation between active layer depth and mean winter air temperature (1 June–31 August in the previous winter) for each site was poor \( R \) values of -0.40 to 0.18) except Minna Bluff \( R = 0.78 \). The poor relationship between mean winter air temperature and active layer depth suggested that the between summer variability in active layer depth at each individual site, except Minna Bluff, responded more to the summer air temperature than to the winter air temperature.

Solar radiation

The correlation between active layer depth and total summer solar radiation (1 December–31 January) at each site was poor \( R \) values of -0.5 to +0.2). The absorption or loss of solar radiation at the soil surface, and thus the amount of heating for the soil profile is expected to be influenced by the surface albedo (Balks et al. 1995, Campbell et al. 1997, 1998). However, the lower albedo sites (Scott Base, Minna Bluff, and Mount Fleming) did not have any improved relationship between solar radiation and active layer depth. Conovitz et al. (2006) also reported no statistically significant relationship between active layer depth and total incoming solar radiation.

**Prediction of active layer depth**

Zhang et al. (1997) showed that there can be no single variable that explains the spatial differences in parameters such as the active layer depth. For example, the mean summer air temperature is not an independent variable as it is affected by other factors such as solar radiation and wind speed. Therefore, a single factor regression of active layer depth and any climatic variable, such as mean summer air temperature, is not likely to fully explain the interannual variation in active layer depth. A step-wise multiple regression analysis with all sites (except Mount Fleming and Granite Harbour, due to missing data) pooled together gave the equation:

\[
\text{ALD} = 222 + 5.69(\text{MSAT}) + 3.63(\text{MWAT}) - 10.6(\text{TSSR}) - 2.84(\text{MSWS})
\]

Where ALD = active layer depth, MSAT = mean summer (December-January) air temperature, MWAT = mean winter (June–August) air temperature, TSSR = total summer solar radiation and MSWS = mean summer wind speed.

Mean summer air temperature, mean winter air temperature, total summer solar radiation and mean summer wind speed together explained 73% of the variation in the active layer depth across all the sites together. The regression had a residual standard deviation of 7.32, and was based on 34 observations across six sites. All four predictors made a significant contribution to the prediction of the active layer depth \( P < 0.02 \). Mean summer air temperature alone explained 15% of the variation in active layer depth. When mean winter air temperature was added as a predictor variable (with mean summer air temperature), 65% of variation in active layer depth could be explained. The \( R^2 \) was further improved to 0.68 by including total summer solar radiation as a predictor variable. Adding the fourth predictor variable, mean summer wind speed improved the \( R^2 \) to 0.73. The regression equation may potentially be used as a predictor for active layer depth at sites in the McMurdo Sound region.
Limitations of the study

The data available form a considerable, and increasingly valuable, database with good supporting soil and site description metadata. However, the interpretation of the data to determine longer-term trends is limited by the relatively short record thus far (eight years at four sites, four years at two sites and three years at one site). The data has further limitations due to incomplete records because of equipment malfunction, particularly in the winter months, with only one full year of data available for Mount Fleming. Also interpretation of the data has to be approached with caution due to some site-specific conditions, such as the occurrence of meltwater flow at Granite Harbour. However, these data do represent one of the most comprehensive soil climate datasets available in the Antarctic, extending through the active layer and into the top part of the permafrost.

Conclusions

The maximum active layer depth was calculated for eight consecutive summers (starting 1999/2000 and ending 2006/07) at seven sites located in the McMurdo Sound region. There was marked interannual variability in both the timing and depth of maximum thaw which was strongly influenced by between-summer variability in mean summer air temperatures. As expected, active layer depth showed clear latitudinal and altitudinal gradients. The local topography had a notable effect at Granite Harbour, where meltwater probably transferred heat to the subsurface and accelerated the thawing process.

Within each site, except for Minna Bluff and Mount Fleming, the active layer increased with increasing mean summer air temperature ($R > 0.87$). At individual sites the between-summer differences in active layer depth did not correlate with mean winter air temperature except at Minna Bluff, nor did active layer depth correlate with total summer incoming solar radiation. Minna Bluff, while not strongly influenced by mean summer air temperature did exhibit an increase in active layer depth with both mean summer wind speed ($R = 0.9$) and mean winter temperature ($R = 0.78$).

When all sites (except Mount Fleming and Granite Harbour) were included together in multiple regression analysis, active layer depth was significantly influenced by mean summer air temperature, mean winter air temperature, total summer solar radiation and mean summer wind speed. When combined, the four predictors accounted for 73% of the variation in the active layer depth, providing a potential mechanism for prediction of active layer depth at other sites in the Ross Sea Region.

Acknowledgements

This research was partially funded by The New Zealand Foundation for Research Science and Technology, (FRST grant CO9X0307) and logistic support was provided by Antarctica New Zealand. Thanks to John Kimble, Ron Paetzold, Don Huffman, Iain Campbell, Jackie Aislabie and Deb Harms who have all assisted with initial climate station establishment and maintenance. Thanks also to Professor Ray Littler of the Statistics Department (University of Waikato) for assistance with statistical advice and analysis. Funding assistance for Leah Adlam from Education New Zealand facilitated collaboration on this paper between the University of Waikato and the USDA-NRCS in Lincoln, Nebraska. Thanks to the referees for helpful comments that led to improvements in the final paper.

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